

Class overview today - November 12, 2018

• Questions about Exercise 2?

• Lecture: **Natural diffusion**

- **•** Introduction to the diffusion process
- Mathematical description of diffusion
- Hillslope diffusion processes

• Exercise 3: **Hillslope diffusion**

Introduction to Quantitative Geology

Natural diffusion: Hillslope sediment transport

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12.11.2018

• Introduce the **diffusion process**

• Present some examples of **hillslope diffusive processes** (heave/creep, solifluction, rain splash)

Diffusion as a geological process

Grain boundary sliding

Rock rheology

Hillslope erosion

156 D.L. Shuster et al. / Earth and Planetary Science Letters 249 (2006) 148–161 4He diffusion in apatite

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General concepts of diffusion

• **Diffusion** is a process resulting in mass transport or mixing as a result of the random motion of diffusing particles

The diffusion process

The diffusion process

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General concepts of diffusion

- **Diffusion** is a process resulting in mass transport or mixing as a result of the random motion of diffusing particles
	- Net motion of mass or transfer of energy is from regions of high concentration to regions of low concentration
	- Diffusion reduces concentration gradients

A more quantitative definition

• **Diffusion** occurs when a **conservative property** moves through space at a **rate proportional to a gradient**

- **Conservative property:** A quantity that must be conserved in the system (e.g., mass, energy, momentum)
- **Rate proportional to a gradient**: Movement occurs in direct relationship to the change in concentration
	- Consider a one hot piece of metal that is put in contact with a cold piece of metal. Along the interface the change in temperature will be most rapid when the temperature difference is largest

- We can now translate the concept of diffusion into mathematical terms.
	- We've just seen "**Diffusion** occurs when a (1) **conservative property** moves through space at a (2) **rate proportional to a gradient**"
- If we start with part 2, we can say in comfortable terms that [transportation rate] is proportional to [change in concentration over some distance]

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- Finally, in symbols we can say

$$
q\ \propto \frac{\Delta C}{\Delta x}
$$

where q is the mass flux, α is the "proportional to" symbol, Δ indicates a

change in the symbol that follows, C is the concentration and x is distance

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- If transport is directly proportional to the gradient, we can replace the proportional to symbol with a constant
- We can also replace the finite changes Δ with infinitesimal changes ∂
- Keeping the same colour scheme, we see

where *D* is a constant called the **diffusion coefficient** or **diffusivity**

- Consider the example to the left of the concentration of some atoms A and B
- Here, we can formulate the diffusion of atoms of A across the red line with time as

$$
q = -D \frac{\partial C_A}{\partial x}
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where C_A is the **concentration** of atoms of A

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- We can consider a simple case for finite changes at two points: (x_1, C_1) and (x_2, C_2)
- At those points, we could say

$$
q = -D\frac{\Delta C}{\Delta x}
$$

$$
q = -D\frac{C_2 - C_1}{x_2 - x_1}
$$

As you can see, ΔC will be negative while Δx is positive, resulting in a negative gradient

Multiplying the negative gradient by $-D$ yields a positive flux q along the x axis, which is what we expect

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 Conservation of mass/energy

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\frac{\Delta C}{\Delta t} = -\frac{q_2 - q_1}{x_2 - x_1}
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- Consider the fluxes q_1 and q_2 at two points, x_1 and x_2
	- What happens when the flux of mass q_2 at x_2 is larger than the flux q_1 at x_1 ?

If we again replace the finite changes Δ with infinitesimal changes ∂ , we can describe our example on the left

$$
\frac{\partial C_{\mathbf{A}}}{\partial t} = -\frac{\partial q}{\partial x}
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Essentially, all this says is that the concentration of A will change based on the flux across a reference face at position x minus the flux across a reference face at position $x + dx$

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- On this week's lesson page you can find notes on how to mathematically combine the two equations we've just seen into the diffusion equation, and how the diffusion equation can be solved
	- Solving the diffusion equation

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General concepts of diffusion

- So our definitions of diffusion to this point are OK for true diffusion processes, but there are also numerous geological processes that are not themselves diffusion processes, but result in diffusion-like behaviour
	- **Hillslope diffusion** is a name given to the overall behaviour of various surface processes that transfer mass on hillslopes in a diffusion-like manner

Erosional processes

• **Erosional processes** are divided between short range (e.g., hillslope) and long range (e.g., fluvial) transport processes

- **Hillslope processes** comprise the different types of mass movements that occur on hillslopes
	- **Slides** refer to **cohesive blocks** of material moving on a well-defined surface of sliding
	- **Flows** move entirely by differential shearing within the transported mass with no clear plane at the base of the flow
	- **Heave** results from disrupting forces acting perpendicular to the ground surface by expansion of the material

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Our focus

• **Heave** results from disrupting forces acting perpendicular to the ground surface by expansion of the material

• Creep is almost too slow

- **Creep: The extremely slow movement of material in response** to gravity
	- **Heave**: The vertical movement of unconsolidated particles in response to expansion and contraction, resulting in a net downslope movement on even the slightest slopes
	- **Seasonal creep** or **soil creep** is periodically aided by heaving

Heave and creep

Nearly vertical Romney shale displaced by seasonal creep

Fig. 4.28, Ritter et al., 2002

Fig. 4.29, Ritter et al., 2002

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How does heaving work?

Fig. 4.30, Ritter et al., 2002

- Near-surface material moves perpendicular to the surface during expansion (E)
	- Expansion can result from swelling or freezing
- In theory, particles settle vertically downward during contraction (C)
- In reality, particle settling is not vertical, but follows a path closer to D

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Based on this concept, what do you think will influence the rates of creep? Slope angle, soil/regolith moisture, particle size/ composition

Common features of hillslope diffusion

- The rate of transport is strongly dependent on the hillslope angle
	- Steeper slopes result in faster downslope transport
	- In other words, the flux of mass is proportional to the topographic gradient
- This suggests these erosional processes can be modelled as **diffusive**

• **What are the two components of diffusion processes?**

• **How does soil creep result in diffusion of soil or regolith?**

• What are the main factors controlling the rate of hillslope diffusion?

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Additional examples of hillslope diffusion

• **Solifluction**

• **Rain splash**

• Tree throw

• Gopher holes

Fig. 11.14b, Ritter et al., 2002

- **Solifluction** occurs in saturated soils, often in periglacial regions
	- In periglacial settings, **frost heave** leads to expansion of the near-surface material
	- During warm periods, saturated material at the surface flows downslope above the impermeable permafrost beneath

[http://geofaculty.uwyo.edu/neil/](http://geofaculty.uwyo.edu/neil/teaching/geologypics/rain%20erosion%20progression.png)

• **Rain splash transport** refers to the downslope drift of grains on a sloped surface as a result of displacement by raindrop impacts

Although this process can produce significant downslope mass transport, it is generally less significant than heave

Studying rain splash

Experimental setup:

- "Rain drops" released from a syringe \sim 5 m above a dry sand target
- Drops travel down a pipe to avoid interference by wind
- Various drop sizes (2-4 mm), sand grain sizes (0.18 - 0.84 mm) and hillslope angles
- High-speed camera used to capture raindrop impact and sand grain motion

small grain ridge in front of a spreading drop front at approximately of the ridge in the ridge is a state in the ridge initial impact of the ridge is a state in the r
The ridge is a state in the ridge is a state in the ridge is a state in the ridge in the ridge in the ridge is UNIVERSITT OF HELSINNI

Furbish et al., 2007

Studying rain splash

• Dry sand grains are displaced following raindrop impact

Miniature bolide impacts (?)

UNIVERSITY OF HELSINKI **Example and and and and bottom part of Cuantitative Geology** and any www.helsinki.fi/yliopisto Furbish et al., 2007 **HELSINGFORS UNIVERSITET** grain collisions; note in the corona-like splash of water (the original motion corona-like splash of water α

Studying rain splash \blacksquare

Figure 11. Radial plots of final ejected grain positions showing increasing asymmetry with increasing Furbish et al., 2007

drift

Biogenic transport: Tree throw

- Falling trees also displace sediment/soil and can produce downslope motion
	- When trees fall, its root mass rotates soil and rock upward
	- Gradually, this soil/rock falls down beneath the root mass as it decays

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Gabet et al., 2003

Biogenic transport: Gopher holes

- Gophers dig underground tunnels parallel to the surface and displace sediment both under and above ground
- On slopes, this sediment is displaced downslope, resulting in mass movement
- Locally, this process can be the dominant mechanism For sediment transport

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